Guervilly C, Hughes DW.

[Jets and large-scale vortices in rotating Rayleigh-Bénard convection.](#)

Physical Review Fluids 2017, 2, 113503.

Copyright:

©2017 American Physical Society

DOI link to article:

<https://doi.org/10.1103/PhysRevFluids.2.113503>

Date deposited:

30/11/2017

Jets and large-scale vortices in rotating Rayleigh-Bénard convection

Céline Guervilly^{1,*} and David W. Hughes²¹*School of Mathematics, Statistics and Physics, Newcastle University,
Newcastle Upon Tyne, NE1 7RU, United Kingdom*²*Department of Applied Mathematics, University of Leeds, Leeds LS2 9JT, United Kingdom*

(Received 9 August 2017; published 28 November 2017)

One of the most prominent dynamical features of turbulent, rapidly rotating convection is the formation of large-scale coherent structures, driven by Reynolds stresses resulting from the small-scale convective flows. In spherical geometry, such structures consist of intense zonal flows that are invariant along the rotation axis. In planar geometry, long-lived, depth-invariant structures also form at large scales, but, in the absence of horizontal anisotropy, they consist of vortices that grow to the domain size. In this work, through the introduction of horizontal anisotropy into a numerical model of planar rotating convection by the adoption of unequal horizontal box sizes (i.e., $L_x \leq L_y$, where the xy plane is horizontal), we investigate whether unidirectional flows and large-scale vortices can coexist. We find that only a small degree of anisotropy is required to bring about a transition from dynamics dominated by persistent large-scale vortices to dynamics dominated by persistent unidirectional flows parallel to the shortest horizontal direction. When the anisotropy is sufficiently large, the unidirectional flow consists of multiple jets, generated on a time scale smaller than a global viscous time scale, thus signifying that the upscale energy transfer does not spontaneously feed the largest available mode in the system. That said, the multiple jets merge on much longer time scales. Large-scale vortices of size comparable with L_x systematically form in the flanks of the jets and can be persistent or intermittent. This indicates that large-scale vortices, either coexisting with jets or not, are a robust dynamical feature of planar rotating convection.

DOI: [10.1103/PhysRevFluids.2.113503](https://doi.org/10.1103/PhysRevFluids.2.113503)

I. INTRODUCTION

In the liquid core of planets, heat is transported primarily by convection. Much effort has been dedicated to the study of the properties of convection, in particular because of its crucial implications for the thermal evolution of planets and the generation of magnetic fields [1,2]. Convection is strongly affected by the rapid rotation of the planet via the action of the Coriolis force, which tends to inhibit small-scale motions along the direction of the rotation axis. Owing to the very low fluid viscosity, flows in planetary cores are turbulent, although the nonlinear inertial effects are relatively weak compared with the Coriolis force. Rotationally constrained convection in planetary cores is thus characterized by small Ekman numbers ($\text{Ek} = \nu/2\Omega d^2$, where ν is the fluid kinematic viscosity, Ω the rotation rate, and d the size of the system), large Reynolds numbers ($\text{Re} = U d/\nu$, where U is a typical flow velocity at the length scale d), and small Rossby numbers ($\text{Ro} = \text{Ek Re} = U/2\Omega d$). Under these conditions, the convective flows take the form of tall narrow columns aligned with the rotation axis. The concerted action of the Reynolds stresses resulting from the small-scale convective columns drives large-scale coherent structures, which can become prominent dynamical features. The occurrence of these large-scale structures in experimental and numerical models of turbulent flows [3,4] attracts much interest because they are observed in many geophysical and astrophysical bodies [5,6]. Although the large-scale flows are usually directed perpendicularly to the direction of gravity, they can nonetheless affect the convective heat transport by shearing the convective cells or

*celine.guervilly@newcastle.ac.uk

by modifying the local properties of the convection [7,8]. Understanding the formation and long-term evolution of the large-scale flows is therefore an essential part of the description of a convective system. Furthermore, large-scale flows can promote the generation of coherent magnetic fields at the system size, while the underlying small-scale turbulent flows may, of themselves, be able to generate only fluctuating small-scale magnetic fields [9,10].

In spherical rotating convection, the large-scale flows consist of intense azimuthal and axisymmetric (i.e., zonal) flows that are invariant along the rotation axis [11]. The presence of the curved spherical boundaries is crucial for the production of such zonal flows [12]. In planar geometry in a horizontally periodic domain, long-lived, depth-invariant structures also form at large scales when the horizontal box sizes are equal. However, the absence of horizontal anisotropy in this case ensures that no persistent unidirectional flow emerges spontaneously; the large-scale structures therefore consist of vortices that grow to the box size [13]. These large-scale vortices (LSVs) form whether the rotation axis is aligned or inclined with respect to the direction of gravity [14]. In this work, we are interested in the possible relation between unidirectional flows and LSVs and whether these two types of large-scale coherent structures can coexist. In simulations of rotating spherical convection, the formation of tall vortices at scales larger than the typical convective size has not been observed. In a Cartesian domain, the size of the LSVs is limited only by the horizontal extent of the periodic domain. If the upscale energy transfer were allowed to continue to larger scales, the LSV would eventually feel either the latitudinal variation of the Coriolis parameter (the β -effect) or the slope of the vertical boundaries. The absence of LSVs in spherical models suggests that the large-scale dynamics prefers unidirectional flows rather than vortices when possible.

The convective flows in a turbulent rotating system are three-dimensional (3D), but strongly anisotropic owing to the rapid rotation. Results from two-dimensional (2D) turbulence might thus be helpful in explaining or predicting some aspects of the dynamics of convective systems on time scales longer than the rotation period when $Ek \ll 1$, $Ro \ll 1$. In particular, the LSVs obtained in 3D convective systems bear a strong resemblance to the condensates that form in 2D turbulence from an upscale kinetic energy transfer [15]. However, the analogy with 2D turbulence remains only partial in this particular case because the convective flows that feed energy into the LSVs are anisotropic but certainly not depth-invariant [7]. The idea that unidirectional flows are preferred to vortices on the large scales once the horizontal symmetry is broken is supported by studies of forced 2D turbulence in which the gradual increase of the β effect leads first to the weakening of the coherent vortices, and eventually to their disappearance, while unidirectional flows emerge [16]. However, a counterexample to the weakening of vortices in the presence of symmetry breaking is demonstrated by the recent study of Frishman *et al.* [17], who considered a forced 2D turbulence model with anisotropy introduced by using a rectangular (nonsquare) periodic domain; with this geometry, they found that both unidirectional flows and vortices emerge at large scales and coexist on long time scales. Here we study whether this remarkable result can carry over to three-dimensional flows.

The two main objectives of the present paper are to determine whether the LSVs that form in rotating planar convection can persist when the symmetry between the horizontal directions is broken and also to determine whether unidirectional flows and LSVs can coexist on long time scales. There are a number of ways of breaking the horizontal symmetry. Considering convection with a tilted rotation axis, thereby modeling midlatitudes, is an obvious physically motivated approach. Here though, we choose to introduce a distinction between the horizontal directions in a three-dimensional Cartesian model of rotating Boussinesq convection by considering rectangular, as opposed to square, horizontal domains. Horizontal anisotropy is thus imposed on the system by allowing larger length scales to exist in one horizontal direction than the other. This formulation is computationally straightforward to implement, thus allowing a thorough exploration of parameter space, and also allows us to make comparisons with the forced 2D turbulence study of Frishman *et al.* [17].

The layout of the paper is as follows. The mathematical formulation of the problem is briefly described in Sec. II. The formation of unidirectional flows and LSVs is described in Sec. III A. The box aspect ratio required for the transition from LSVs to unidirectional flows is quantified in

Sec. III B, and the selection of the length scale of the large-scale flow is described in Sec. III C. The occurrence of bistable states is presented in Sec. III D. In Sec. III E, we establish the domain of existence of the large-scale flows. Finally, in Sec. III F, we discuss the effect of the large-scale flows on the convective heat transfer. A conclusion is given in Sec. IV.

II. MATHEMATICAL FORMULATION

We use a local planar model of rotating Boussinesq convection, as in Ref. [7]. The computational domain is three-dimensional and periodic in the horizontal directions. A vertical temperature difference, ΔT , is imposed across the layer of depth d . In the horizontal plane, the size of the domain is $L_x d$ in the x direction and $L_y d$ in the y direction. The gravitational field is uniform, $\mathbf{g} = -g\mathbf{e}_z$. The rotation vector is $\Omega\mathbf{e}_z$. The fluid has kinematic viscosity ν , thermal diffusivity κ , density ρ , and thermal expansion coefficient α , all of which are constant. Lengths are scaled with d , times with $1/(2\Omega)$ and temperature with ΔT . The system of dimensionless governing equations thus becomes

$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} + \mathbf{e}_z \times \mathbf{u} = -\nabla p + \text{Ek} \nabla^2 \mathbf{u} + \frac{\text{Ra} \text{Ek}^2}{\text{Pr}} \theta \mathbf{e}_z, \quad (1)$$

$$\nabla \cdot \mathbf{u} = 0, \quad (2)$$

$$\frac{\partial \theta}{\partial t} + \mathbf{u} \cdot \nabla \theta - u_z = \frac{\text{Ek}}{\text{Pr}} \nabla^2 \theta, \quad (3)$$

where $\mathbf{u} = (u_x, u_y, u_z)$ is the velocity field, p the pressure, and θ the temperature perturbation relative to a linear background profile. The dimensionless parameters are the Rayleigh number, $\text{Ra} = \alpha g \Delta T d^3 / \kappa \nu$, the Ekman number, $\text{Ek} = \nu / 2\Omega d^2$, and the Prandtl number, $\text{Pr} = \nu / \kappa$. The upper and lower boundaries are taken to be perfect thermal conductors, impermeable and stress free. Stress-free boundary conditions, where boundary friction is absent, favor the emergence of large-scale flows of large amplitude in both planar [18,19] and spherical geometry [8]. Equations (1)–(3) are solved using the pseudospectral code described in detail in Ref. [20].

III. RESULTS

A. Jets and large-scale vortices

For most of the results discussed in the paper, the parameters are fixed as $\text{Ek} = 10^{-5}$, $\text{Ra} = 3 \times 10^8$, $\text{Pr} = 1$, $L_x = 1$, with variations only in L_y . The Ekman and Rayleigh numbers are varied only in Secs. III C and III D, in order to integrate the system over many viscous time scales, and in Sec. III E, in order to determine the domain of existence of the large-scale flows; furthermore, we perform one run with $L_x = 2$ and $L_y = 4$, which we discuss in the conclusion. The value of the Rayleigh number $\text{Ra} = 3 \times 10^8$ corresponds to 7.4Ra_c , where Ra_c is the critical Rayleigh number at the linear onset of convection [21].

Table I contains some global quantities obtained from simulations with L_y in the range $1 \leq L_y \leq 8$, together with the numerical resolutions used. The root mean square (r.m.s.) amplitude of the velocity is measured by the Reynolds number $\text{Re} = \langle |\mathbf{u}| \rangle^{1/2} \text{Ek}^{-1}$, and the r.m.s. amplitude of the i component of the velocity by $\text{Re}_i = \langle u_i^2 \rangle^{1/2} \text{Ek}^{-1}$, where the angle brackets denote volume and time averages. The inclusion of the Ekman number in these definitions is due to our choice of dimensional units. In order to confirm that the resolution was adequate, and also to verify that the results were not critically dependent on the number of spectral modes adopted, provided this was sufficiently large, we simulated the case of $L_y = 3$ with both 512 and 768 modes in the y direction. As can be seen from Table I, the differences between the two simulations are small.

TABLE I. Simulations run at varying L_y for $Ek = 10^{-5}$, $Ra = 3 \times 10^8$, and $L_x = 1$. The value n denotes the number of positive (or, equivalently, negative) jets of the mean velocity u_x . The quantities Γ and α are defined in Sec. IIIB. The final column lists the numerical resolutions employed, with N_i indicating the number of collocation points in the direction i .

L_y	Re_x	Re_y	Re_z	Re	Nu	Γ	α	n	$N_x \times N_y \times N_z$
1	2469	2474	910	3612	28.2	5.3	0.00	0	$256 \times 256 \times 129$
1.02	2589	2443	917	3676	28.4	5.4	0.06	0	$256 \times 256 \times 129$
1.05	2811	2298	915	3744	28.4	5.6	0.20	0	$256 \times 256 \times 129$
1.08	3130	880	922	3380	29.1	4.5	0.85	1	$256 \times 256 \times 129$
1.1	3181	874	925	3427	29.3	4.6	0.86	1	$256 \times 256 \times 129$
1.5	4177	887	907	4365	28.6	7.7	0.91	1	$256 \times 256 \times 129$
2	5548	956	910	5702	28.6	13.1	0.94	1	$256 \times 512 \times 129$
2.1	5612	1120	906	5792	28.5	12.8	0.92	1	$256 \times 512 \times 129$
2.5	6406	1073	901	6555	28.3	17.6	0.95	1	$256 \times 512 \times 129$
3	4393	924	907	4579	28.5	8.5	0.92	2	$256 \times 512 \times 129$
3	4465	925	911	4647	28.7	8.7	0.92	2	$256 \times 768 \times 129$
4	5621	1002	894	5779	27.6	13.9	0.94	2	$256 \times 512 \times 129$
6	4505	966	909	4697	28.5	8.9	0.91	4	$256 \times 1024 \times 129$
8	4681	943	900	4856	27.9	9.7	0.92	5	$256 \times 1024 \times 129$

1. Initial growth of the convective and large-scale flows

Unless stated otherwise, the simulations described were started from a small perturbation to the basic state of rest. Figure 1 shows the time series of $\hat{Re}_x$, $\hat{Re}_y$, and $\hat{Re}_z$, which are defined as the Reynolds number Re_i except with only a spatial average, for $L_y = 8$. The figure shows the evolution only for 12 000 dimensionless time units, i.e., 0.12 of the global viscous time scale; the full simulation was actually run for one viscous time scale. We first describe the initial growing phase of the kinetic energy; the saturated phase is detailed in the following subsections. Initially, the

FIG. 1. Time series of the r.m.s. values of the three velocity components for $L_y = 8$. The gray dashed lines indicate the times of the snapshots in Fig. 2.

FIG. 2. Snapshots of the depth-averaged axial vorticity taken at the times indicated by the gray dashed lines in Fig. 1.

kinetic energy grows exponentially, with the velocity characterized by small horizontal scales. This short period of rapid growth in the kinetic energy is followed by an abrupt fall once the nonlinear terms become important. At this stage the convection remains small scale with a horizontal scale of approximately 0.1 for the chosen parameters. Subsequently the kinetic energy grows again, this time on a much longer time scale, as the flow moves to large horizontal scales.

Figure 2 shows snapshots of the depth-averaged axial vorticity (denoted by $\langle \omega_z \rangle_z$, where $\langle \cdot \rangle_i$ indicates an average along the direction i) at the different times indicated in Fig. 1. This representation focuses on the formation of large-scale depth-invariant flows; the convective flows, whose axial vorticity is mainly antisymmetric with respect to the midplane [21], are not visible on these plots. During the period of slow evolution to large-scale flows, vortices of large horizontal scale compared with the convective scale first become established across the domain [Fig. 2(a)], with movement of the vortices in both horizontal directions. The growth of these vortices corresponds to an increase in both $\hat{Re}_x$ and $\hat{Re}_y$, with $\hat{Re}_z$ saturating at a lower value; these large-scale vortices consist essentially of depth-invariant horizontal flows. Vortex mergers occur and the LSVs grow to a scale comparable with the shortest horizontal box size [Fig. 2(b)]. Subsequently, a number of jets become established, parallel to the shortest horizontal direction [Fig. 2(c)]. The LSVs persist in the presence of the jets and, in the saturated state, move only with the jets. The growth of kinetic energy is eventually saturated on a time scale that is long compared with the dynamical time scale of the convection but short compared with the viscous time scale.

2. Flow in the saturated phase for $L_y = 1$

We now describe the flow in the final saturated state, first recalling results obtained for the standard case of $L_y = 1$. Figure 3 shows snapshots of horizontal and vertical cross sections of ω_z and u_z . The dominant feature in the plot of ω_z is a large-scale vortex that is predominantly z invariant. The LSV consists of a concentrated cyclone and a more dilute anticyclone, which grow until they reach the box size. LSVs are a robust feature of turbulent, rotationally constrained convection in Cartesian

FIG. 3. Snapshots of the horizontal and vertical cross sections of (a) the axial vorticity and (b) the vertical velocity for $L_y = 1$.

geometry when $L_x = L_y$ [7,13,15,18,22,23]. The asymmetry between cyclone and anticyclone is thought to be due both to the instability of the large-scale anticyclone, whose local vorticity is comparable with the planetary vorticity 2Ω , and to the preference for cyclonic convective vortices [7]. This asymmetry eventually disappears as the Rossby number is lowered via a decrease of the Ekman number [18,24].

In spectral space, the LSV is a superposition of the wave numbers $(k_x, k_y) = (0, 1)$ and $(k_x, k_y) = (1, 1)$ for the velocity component u_x and of $(k_x, k_y) = (1, 0)$ and $(k_x, k_y) = (1, 1)$ for u_y . The LSV therefore has a visible signature in snapshots of $\langle u_x \rangle_{xz}$ and $\langle u_y \rangle_{yz}$. However, the LSV drifts in random horizontal directions over time, so long time averages of either $\langle u_x \rangle_{xz}$ or $\langle u_y \rangle_{yz}$ produce a zero profile. In this sense, the LSVs do not produce any persistent unidirectional flow.

The u_z plot of Fig. 3 shows that small-scale convective flows are embedded within the large-scale cyclone. The convective structures can be elongated in one horizontal direction, especially in the regions of intense shear surrounding the LSV. There is no discernable box-size vertical flow associated with the presence of the LSV, so LSVs do not directly contribute to the outward heat transport, although they might have a feedback effect [7,24].

3. Flow in the saturated phase for $L_y > 1$

Figure 4 shows snapshots of the horizontal and vertical cross sections of ω_z and u_z for $L_y = 8$ in the final saturated state. The dominant feature in the plot of ω_z is the unidirectional flow that consists of five jets, each with a cyclonic and anticyclonic component, directed along x . A multitude of small-scale convective vortices are embedded within the jets. The cyclonic LSVs that are prominent in the plots of $\langle \omega_z \rangle_z$ during the initial growing phase shown in Fig. 2 are also visible in the saturated phase. However, as seen in Fig. 2(c), they are no longer located in the middle of the bands of mean (i.e., x -averaged) cyclonic axial vorticity, but lie towards the flanks. The plot of u_z shows that the convective structures can be elongated along the x direction in the shear regions but that there is no discernible box-size vertical flow associated with the presence of the jets. Figure 5 shows snapshots of $\langle \omega_z \rangle_z$ and of the profile of $\langle u_x \rangle_{xz}$ in the saturated stage for $L_y = 8$ and for the intermediate cases $L_y = 2$ and 4. The unidirectional flow consists of one jet for $L_y = 2$ and two jets for $L_y = 4$. No persistent unidirectional flow along y is produced.

The presence of a unidirectional flow might be expected in this system by analogy with 2D turbulence, in which the upscale energy transfer feeds the largest mode available [25]. However, as

FIG. 4. Snapshots of the horizontal and vertical cross sections of (a) the axial vorticity and (b) the vertical velocity for $L_y = 8$.

noted in Ref. [17] in a study of forced 2D turbulence, the largest mode argument does not explain why there is more than one stable jet for $L_y = 4$ and $L_y = 8$. In all cases studied, LSVs coexist with the jets. The LSVs are advected by the mean flow u_x and always remain at the same location in y . At most two cyclones can coexist in the same cyclonic band, and they remain in the flanks of their respective jets, with no cyclone sharing its y location. The cyclones are never observed at the location corresponding to $\langle u_x \rangle_{xz} = 0$, which is the region of maximum shear $\partial_y \langle u_x \rangle_{xz}$. Furthermore, no large-scale anticyclone is visible in the bands of anticyclonic vorticity; hence the asymmetry between large-scale cyclones and anticyclones is present for $L_y > L_x$, as it is when $L_y = L_x$. The mean flows themselves, however, are not asymmetric, since cyclonic and anticyclonic bands have comparable width (and hence comparable amplitude).

B. How much horizontal anisotropy is required to drive the jets?

Since large-scale motions consist mostly of horizontal flows, we follow Ref. [7] in using the parameter $\Gamma = \text{Re}^2 / (3\text{Re}_z^2)$ to determine their presence; $\Gamma > 1$ indicates the presence of either jets or LSVs. Such is the case for all the simulations reported in Table I. To distinguish between those simulations in which only LSVs form and those in which jets are present, we calculate the coefficient α defined by

$$\alpha = \frac{\langle u_x^2 \rangle - \langle u_y^2 \rangle}{\langle u_x^2 \rangle + \langle u_y^2 \rangle}, \quad (4)$$

FIG. 5. Snapshots of the depth-averaged axial vorticity (bottom panel) and profile of the mean u_x as a function of y (top panel) for (a) $L_y = 2$, (b) $L_y = 4$, and (c) $L_y = 8$.

which measures the horizontal anisotropy of the flow; α is close to unity when jets are present. Figure 6 shows α as a function of L_y ; crucially, only a small deviation of L_y from L_x (namely, $L_y = 1.08$ for the parameters used here) is required to observe the spontaneous emergence of jets.

FIG. 6. Flow anisotropy coefficient α , defined by (4), as a function of L_y .

FIG. 7. Space-time diagram of the depth-averaged axial vorticity for $\text{Ek} = 10^{-5}$, $\text{Ra} = 3 \times 10^8$, $L_x = 1$, and $L_y = 4$: (a) x -averaged and (b) the r.m.s. of the fluctuating component.

C. Long-term evolution and merging of the jets

To study the drifting or merging of the jets and LSVs over time, we plot in Fig. 7 the space-time diagram of the axial vorticity averaged in x and z , $\langle \omega_z \rangle_{xz}$ (the mean vorticity), together with the r.m.s. of the fluctuating component, defined as $\langle \omega_z \rangle_z - \langle \omega_z \rangle_{xz}$. The mean vorticity contains information about both the jets and the LSV, while the fluctuating part contains information about the LSV. Over this time series, the jets do not drift or merge; indeed, this is representative of all the cases of Table I, irrespective of the number of jets. The absence of measurable drift is consistent with the fact that the horizontal average of the horizontal velocity is negligible in these simulations. LSVs can persist in the same band of mean cyclonic vorticity for the whole duration of the time integration or they might be intermittent. LSVs on opposite flanks of the same cyclonic band are advected in opposite directions and do not appear to interact with each other. However, in some occurrences, the formation of a new LSV might destabilize the neighboring LSV, since the disappearance of an LSV sometimes closely follows the appearance of its immediate neighbor. It is plausible that the persistence or intermittency of the LSVs is linked to the width of the band of mean vorticity in which they are located. Indeed, in Fig. 7, the LSVs within the narrower band located around $y = 1$ are more intermittent than those within the wider band located around $y = 3$. The comparison of the behavior of the LSVs for the cases $L_y = 2$ and 2.5 (for which one jet is present) also supports this argument.

Table I gives the number of jets, denoted by n , where, by jets, we mean unidirectional flows that persist in time averages. For clarification, we reiterate that we are defining a jet as consisting of both its cyclonic and anticyclonic component (i.e., both a red and a blue region in plots such as Fig. 2). Note that simulations that produce only LSVs (i.e., $L_y < 1.08$) have $n = 0$; cases with $L_y \geq 3$ have multiple jets ($n > 1$). This result is somewhat surprising since, in a system where there is no

FIG. 8. (a) Time series of the r.m.s. values of the three velocity components and (b) space-time diagram of the depth-averaged axial vorticity for $Ek = 10^{-4}$, $Ra/Ra_c = 4.2$, and $(L_x, L_y) = (1, 4)$.

gradient of the planetary rotation (i.e., no β -effect), one might expect the upscale energy transfer to persist until the largest available scale is attained. Over the entire duration of the simulations with $Ek = 10^{-5}$ there is no evidence of a trend for any slow drift or merger of the jets. That said, for such a small value of Ek , computational limitations restrict the time integration to less than one viscous time scale (which corresponds to Ek^{-1} in our time unit). In all the simulations reported so far, we have described the large-scale structure that gradually emerges from an initial small-scale perturbation and that persists over the entire time integration. The system spontaneously adopts a number of jets such that $L_y/(2n) \approx L_x$ when possible (i.e., when L_y is large enough). This observation can be explained by the snapshots of the initial growing stage in Fig. 2: the LSVs form first and grow to the size L_x , before extending into jets that retain the initial size of the LSV in the y direction. It is, however, plausible that on longer time scales the jets could readjust their size by merging.

In order to investigate very long term aspects of the dynamics, it is necessary to consider an increased value of the Ekman number. To this end, we perform a simulation with $Ek = 10^{-4}$, $Ra = 8 \times 10^6$ ($Ra/Ra_c = 4.2$), and $(L_x, L_y) = (1, 4)$ for a duration of five viscous time scales. The time series of $\hat{Re}_x$, $\hat{Re}_y$, and $\hat{Re}_z$, together with the space-time diagram of $\langle \omega_z \rangle_{xz}$, are shown in Fig. 8. The initial phases of growth and saturation, before one viscous time scale has elapsed, are

FIG. 9. Re_x as a function of L_y/n (for cases with $n > 0$) for $Ek = 10^{-5}$, $Ra = 3 \times 10^8$, and $L_x = 1$.

similar to the case already described for $Ek = 10^{-5}$, with the formation of four jets. Note that no persistent LSV is observed in the jet flanks here because L_x is not large enough to permit it at this value of Ek . Over approximately two viscous time scales, one of the mean cyclonic bands slowly drifts, preceding a relatively rapid expansion and merger, with only one jet remaining; this jet merger is accompanied by an increase by a factor two of $\hat{Re}_x$. On long time scales, the jets therefore evolve by increasing their length scale as a result of the merger of bands of mean cyclonic vorticity, leading to an increase in the amplitude of u_x . This relation between the amplitude of u_x and the jet length scale is consistent with the results presented in Table I (for which $Ek = 10^{-5}$). Figure 9 shows Re_x as a function of the jet length scale L_y/n for the cases with $n > 0$ of Table I, corroborating the linear dependence of Re_x on L_y/n .

D. Bistability of the large-scale flow

In their study of forced 2D turbulence, Bouchet and Simonnet [25] show that when the aspect ratio of the computational domain slightly exceeds unity, the flow switches randomly between a dipole vortex and a unidirectional flow. For the ratio $L_y/L_x = 1.02$, their system spends more time in the dipole vortex configuration, in contrast with the case of $L_y/L_x = 1.04$, for which the unidirectional flow configuration is preferred. These topological changes are found to be slow processes that occur over a viscous time scale, and which therefore can be observed only through long time integrations. In all our simulations reported in Table I (all of which have $Ek = 10^{-5}$), which are started from an initial small-scale perturbation, no random topological changes were observed, even for small deviations of L_y/L_x from unity. However, the results of Ref. [25] show that the solution might be dependent on initial conditions and may vary on viscous time scales.

To study whether the two configurations—LSV-only or jets—can both occur for a given aspect ratio L_y/L_x , we should ideally perform simulations for several viscous time scales. Since computational constraints exclude this possibility for small Ekman numbers, we use two different approaches. First, we force the system to adopt a given large-scale configuration by using an initial condition of finite amplitude. To determine whether the jets can be stable when $L_x = L_y$, we use the data from a snapshot of the case $(L_x, L_y) = (0.5, 1)$ (which is thus in the jet configuration), which we replicate in the x direction to obtain a box of dimension $(L_x, L_y) = (1, 1)$. This fabricated data, with the addition of a small amount of noise, is then used as an initial condition for the domain

FIG. 10. Time series of the r.m.s. values of the three velocity components for the simulation with $(L_x, L_y) = (1, 1)$ whose initial condition is constructed from $(L_x, L_y) = (0.5, 1)$ ($Ek = 10^{-5}$, $Ra = 3 \times 10^8$).

$(L_x, L_y) = (1, 1)$. Figure 10 shows the time series of $\hat{Re}_x$, $\hat{Re}_y$, and $\hat{Re}_z$. The jet configuration is recognizable initially by $\hat{Re}_x > \hat{Re}_y \approx \hat{Re}_z$. After about 1000 time units, or equivalently about 25 turnover time scales based on the horizontal velocity and the box size, the jet configuration is lost in favor of the LSV configuration (where $\hat{Re}_x \approx \hat{Re}_y$). Furthermore, we perform simulations for which a snapshot from a case with $(L_x, L_y) = (1, 1.08)$ (a jet configuration) is used as an initial condition for either $(L_x, L_y) = (1, 1)$ or $(L_x, L_y) = (1.08, 1.08)$. In both cases, the large-scale flow rapidly evolves from jets to an LSV, similarly to the case shown in Fig. 10. Based on these numerical simulations, we conclude that the jets become rapidly unstable when $L_x = L_y$.

The second approach consists, as in Sec. III C, in increasing the Ekman number in order to simulate many viscous time scales. We perform a simulation with $Ek = 10^{-4}$, $Ra/Ra_c = 4.2$, and $(L_x, L_y) = (2, 2.15)$ for a duration of 10 viscous time scales. The time series of $\hat{Re}_x$, $\hat{Re}_y$, and $\hat{Re}_z$ are shown in Fig. 11(a). An LSV forms initially, followed by spontaneous topological changes that occur after the elapse of more than one viscous time scale, which corresponds to 10^4 time units in Fig. 11(a). The LSV configuration is characterized by $\hat{Re}_y > \hat{Re}_z$ and is illustrated in Fig. 11(b), which shows a snapshot of $\langle \omega_z \rangle_z$. The jet configuration is characterized by $\hat{Re}_y \approx \hat{Re}_z$ and is illustrated by Fig. 11(c). For this particular aspect ratio ($L_y/L_x = 1.075$), the system spends most of its time in the LSV configuration, i.e., the probability of obtaining the LSV configuration is higher. As L_y/L_x decreases towards unity, the probability of a jet configuration decreases. Conversely, as L_y/L_x increases, the probability of achieving the LSV configuration decreases. This is illustrated in Fig. 12, which shows the times series of $\hat{Re}_x$, $\hat{Re}_y$, and $\hat{Re}_z$ for $(L_x, L_y) = (2, 2.2)$. We find that for $L_y/L_x = 1.2$, the system remains in the jet configuration during the whole time integration of 10 viscous time scales. For the parameters studied in this paragraph, there is a narrow window, given by $L_y/L_x \in [1.06, 1.15]$, in which both the jet state and the LSV-only state can exist.

E. Domain of existence

When $L_y = L_x$, LSVs form when the convective vortices are constrained by rotation to be anisotropic (i.e., narrow in the horizontal directions and tall in the vertical) and when they are sufficiently energetic to merge [7]. At least for the Ekman numbers considered in Guervilly *et al.* [7]

FIG. 11. (a) Time series of the r.m.s. values of the three velocity components for the simulation with $\text{Ek} = 10^{-4}$, $\text{Ra}/\text{Ra}_c = 4.2$, and $(L_x, L_y) = (2, 2.15)$. (b), (c) Snapshots of the z -averaged axial vorticity at the times indicated by the dashed vertical lines in (a).

($\text{Ek} \geq 5 \times 10^{-6}$), these conditions are quantified by (i) small local Rossby numbers, $\text{Ro}_z^l = \langle u_z \rangle / (2\Omega l) \lesssim 0.1$, where l is the typical horizontal length scale of the convection, and (ii) $\text{Ra}/\text{Ra}_c \gtrsim 3$. To meet both criteria, the Ekman number must be suitably small. An additional condition for the formation of LSVs is that the computational domain must be wide enough to accommodate a sufficient number of convective cells in the horizontal direction (typically more than four convective cells along L_x or L_y).

To determine whether similar constraints apply to the formation of the jets, we first performed simulations at larger Ekman numbers than considered above. Table II lists a number of outputs from simulations run at $\text{Ek} = 3.16 \times 10^{-4}$ and $\text{Ek} = 10^{-4}$ for $\text{Ra}/\text{Ra}_c \approx 4$ and for various box configurations. The ratio Γ is again used to indicate the formation of dominant horizontal flows (related to the formation of large-scale flows) when greater than unity, and the coefficient α is used

FIG. 12. Time series of the r.m.s. values of the three velocity components for the simulation with $\text{Ek} = 10^{-4}$, $\text{Ra}/\text{Ra}_c = 4.2$, and $(L_x, L_y) = (2, 2.2)$.

to indicate the formation of unidirectional flows when close to unity. For $\text{Ek} = 3.16 \times 10^{-4}$ and all the box configurations, the local Rossby number is $\text{Ro}_z^l = 0.12$ and $\Gamma \approx 1$; neither an LSV nor jets form because conditions (i) and (ii) cannot be met simultaneously for this Ekman number. Note that, for this value of Ek , only about three convective cells fit along a unit horizontal length; we thus have to use wide computational domains. When Ek is decreased to 10^{-4} , the local Rossby number decreases to 0.09 for this value of Ra/Ra_c . For all box configurations, $\Gamma > 1$, i.e., large-scale horizontal flows are predominant. For $L_x = L_y = 2$, $\alpha = 0$, i.e. the large-scale flow is an LSV, as expected. For $(L_x, L_y) = (1, 2)$, $\alpha > 0$, reflecting the formation of jets. These observations are all confirmed by visual inspection of the axial vorticity (not shown here for brevity). These simulations thus show that the constraint $\text{Ek} \lesssim 10^{-4}$ is required for the formation of jets, a similar criterion to that for the formation of LSVs.

TABLE II. Simulations performed to determine the domain of existence of the large-scale flows.

Ek	Ra	Ra/Ra_c	L_x	L_y	Re_x	Re_y	Re_z	Re	Ro_z^l	Γ	α	$N_x \times N_y \times N_z$
3.16×10^{-4}	1.7×10^6	4.1	2	2	109	110	119	195	0.12	0.9	0.00	$128 \times 128 \times 97$
3.16×10^{-4}	1.7×10^6	4.1	1	4	119	105	119	198	0.12	0.9	0.12	$64 \times 256 \times 97$
3.16×10^{-4}	1.7×10^6	4.1	2	4	111	109	119	196	0.12	0.9	0.02	$128 \times 256 \times 97$
10^{-4}	8×10^6	4.2	2	2	427	426	202	636	0.091	3.3	0.00	$128 \times 128 \times 97$
10^{-4}	8×10^6	4.2	1	2	445	177	208	523	0.092	2.1	0.73	$64 \times 128 \times 97$
10^{-5}	5×10^7	1.2	1	2	44	43	45	76	0.0045	0.95	0.02	$128 \times 256 \times 129$
10^{-5}	1×10^8	2.5	1	2	389	266	222	521	0.022	1.84	0.36	$128 \times 256 \times 129$
10^{-5}	1.5×10^8	3.7	1	2	2139	438	444	2228	0.041	8.40	0.92	$128 \times 256 \times 129$
10^{-5}	3×10^8	7.4	1	2	5548	956	910	5702	0.084	13.1	0.94	$256 \times 512 \times 129$
10^{-5}	4×10^8	9.9	1	2	6824	1095	1169	7010	0.107	12.0	0.95	$256 \times 512 \times 129$
10^{-5}	6×10^8	14.8	1	2	7708	1365	1551	7980	0.139	8.82	0.94	$256 \times 512 \times 257$
10^{-5}	7×10^8	17.3	1	2	7811	1573	1713	8150	0.153	7.54	0.92	$256 \times 512 \times 257$

FIG. 13. Γ as a function of (a) the normalized Rayleigh number and (b) the local Rossby number for $Ek = 10^{-5}$.

To test whether both of the conditions (i) and (ii) apply to the formation of jets, we perform a number of simulations with increasing Ra for $Ek = 10^{-5}$ and with $(L_x, L_y) = (1, 2)$ (listed in Table II). Figure 13(a) shows the value of Γ as a function of Ra/Ra_c for these cases, which is compared with cases with $(L_x, L_y) = (1, 1)$ at the same Ek as used in Guervilly *et al.* [7]. Γ increases well above unity for both $L_y = 1$ and $L_y = 2$ when $Ra/Ra_c > 3$. Condition (ii) is therefore quantitatively the same for both jets and LSVs. Figure 13(b) shows Γ as a function of Ro_z^l and indicates that condition (i) is quantitatively different for jets ($L_y = 2$) and LSVs ($L_y = 1$): the amplitude of the jets start to decay when the local Rossby number is greater than 0.08, which is smaller than the Rossby number at which Γ starts to decline for the LSV. We thus conclude that the jets are more fragile than the LSVs when the local Rossby number approaches 0.1. Note that for the largest Rayleigh number computed for $L_y = 2$, Γ takes similar values for $L_y = 1$ and $L_y = 2$; however, the large-scale flow for $L_y = 2$ still consists of an $n = 1$ jet accompanied by LSVs in the mean cyclonic region.

F. Effect on the heat transfer

We anticipate that convective heat transfer might be impeded by the presence of jets due to either (i) the shear flow itself or (ii) the local increase of the rotation rate in the regions of large-scale

FIG. 14. Nu as a function of the Reynolds number based on u_x for the simulations of Table I ($L_x = 1$ and varying L_y) and for a reference case without a large-scale flow ($L_x = L_y = 0.3$).

FIG. 15. Profiles of (a) the time-averaged mean flow, $\langle u_x \rangle_{xz}$, (b) the mean axial vorticity, $\langle \omega_z \rangle_{xz}$, and (c) the mean heat flux at the upper boundary, q ; $\text{Ek} = 10^{-5}$, $\text{Ra} = 3 \times 10^8$, $L_x = 1$, and $L_y = 4$. The red bands correspond to the regions of positive mean vorticity.

cyclonic vorticity. It should though be noted that the influence of a mean shear flow on convection is complex and may even enhance the efficiency of the heat transport [26]. Effect (i) is expected to be important when most of the convective heat transport is directed parallel to the mean velocity gradients. Such is the case outside the tangent cylinder in spherical geometry [8], and in planar geometry when the rotation axis is perpendicular to the direction of gravity [27]. In our system, the effect of the shear is difficult to predict: on the one hand, the convective heat flux is perpendicular to the mean velocity gradients, so the shear might not be particularly disruptive; on the other, the elongation of the vertical flows in the x direction observed in Fig. 4(b) indicates that the vertical velocity is affected by the shear. Effect (ii) is due to the well-known inhibiting effect of rotation on convection [21] and has previously been observed in the presence of LSVs when $L_x = L_y$; here the contribution of the LSV vorticity to the background rotation is not negligible locally [7]. The decrease in the heat transport efficiency associated with LSVs depends on the horizontal box size since wider horizontal domains promote the formation of broader LSVs of increasing strength. In the case of jets, the efficiency of the heat transfer may also depend on their amplitude and length scale.

The global efficiency of the heat transfer is often quantified by the Nusselt number, Nu , which is a measure of the total heat flux through the layer normalized by the heat flux in the absence

of convective motions. In our system, in the absence of heat sources and with fixed-temperature boundary conditions, the Nusselt number is defined as

$$\text{Nu} = 1 + \langle |\nabla\theta|^2 \rangle, \quad (5)$$

where the angle brackets denote time and volume averages. To compare quantitatively the effect of the presence of large-scale flows on Nu, we need to calculate Nu in a reference case without large-scale flows. This can be obtained by considering a small horizontal box size that accommodates only about three convective cells, thereby preventing the formation of a large-scale flow. For $\text{Ek} = 10^{-5}$ and $\text{Ra} = 3 \times 10^8$, the reference case is computed for $L_x = L_y = 0.3$. Figure 14 shows the Nusselt number as a function of Re_x for the simulations of Table I and for the reference case. Overall, Nu is smaller by approximately 10% in cases with a large-scale flow than it is in the reference case. As we might have expected, there is a decreasing trend of Nu with Re_x for the cases with jets ($n > 0$). However, at similar values of Re_x , Nu is relatively unaffected by the number of jets. The cases with only an LSV ($n = 0$) have smaller values of Nu than cases with jets with similar values of Re_x .

We now examine the local effect of the large-scale flows on the heat transfer. Figure 15 shows the time-averaged heat flux at the upper boundary, $q = -\langle \partial_z T(z=1) \rangle_x$, for $(L_x, L_y) = (1, 4)$. For comparison, we also plot the time averages of $\langle u_x \rangle_{xz}$ and $\langle \omega_z \rangle_{xz}$. The heat flux is lower by approximately 25% in the regions of positive $\langle \omega_z \rangle_{xz}$ than in the regions of negative $\langle \omega_z \rangle_{xz}$. This main difference is not due to the shearing of the convective flows by the jets because the highs of q are not correlated with regions of smallest shear. The heat flux is particularly low in the locations of the large-scale cyclone, which appear distinctly as peaks in the profile of $\langle \omega_z \rangle_{xz}$. This detrimental effect of large-scale cyclonic flow on the heat transport is due to the local increase of rotation rate, as in the case of the LSVs when $L_x = L_y$. This is consistent with the fact that the mean axial vorticity persistently reaches values of 0.2 (in units of 2Ω). In the reference case $L_x = L_y = 0.3$, $q \approx 18$ on average, so this argument also explains why the anticyclonic regions enhance the heat flux. In cases with LSVs only ($L_y \in [1, 1.05]$), the mean axial vorticity persistently reaches values of 0.5, which explains why LSVs have a more pronounced effect than jets on Nu (Fig. 14).

IV. CONCLUSION

We have studied the formation of large-scale flows in rotating anisotropic convection; this has been achieved through numerical simulations in planar geometry, with horizontal anisotropy introduced by choosing a box aspect ratio with unequal horizontal dimensions ($L_y > L_x$). For the standard case of $L_y/L_x = 1$, which has received much attention in the literature [7, 13, 15, 18, 22–24], the stable configuration consists of a depth-invariant large-scale vortex, formed by the clustering of small-scale convective vortices, and which grows to the horizontal box size. When L_y/L_x is increased above unity, the system undergoes a transition from the LSV-only configuration to one in which depth-invariant unidirectional flows emerge. As for LSVs, the necessary conditions for the formation of these unidirectional flows are small Rossby numbers (based on the convective length scale and velocity) and sufficiently large Reynolds numbers. Over less than one global viscous time scale, persistent unidirectional flows with multiple jets are produced when $L_y \gtrsim 3L_x$. On longer time scales, the multiple jets merge to reach the largest available length scale in the system. Interestingly, we found that LSVs of size comparable with L_x systematically coexist with the jets. The LSVs are located in the flanks of the jets and can be persistent or intermittent depending on the length scale of the jet. The transition from the LSV-only configuration to a configuration involving both jets and LSVs occurs for a small degree of anisotropy, namely, $L_y/L_x \approx 1.08$ when $\text{Ek} = 10^{-5}$ and $L_x = 1$. In the vicinity of this transition, the system is bistable; random changes between the two large-scale states are observed when the simulations are integrated over several viscous time scales. This bistable regime is observed only for a narrow interval of the box aspect ratio ($L_y/L_x \in [1.06, 1.15]$ for $\text{Ek} = 10^{-4}$ and $L_x = 2$).

Large-scale flows are known to influence the efficiency of the heat transfer in rotating convection [8, 27]. In our model, although there are no sizable vertical flows associated with the presence of the

FIG. 16. Snapshot of the depth-averaged axial vorticity in the saturated regime for the case of $L_x = 2$, $L_y = 4$.

LSVs or the jets, we find that the large-scale flows modify the transport properties of the convection by changing locally the rotation rate. As the axial vorticity of the large-scale flows is skewed towards cyclonic vorticity in our simulations, these flows reduce the efficiency of the convective heat transport. In more realistic planetary conditions at smaller Ekman and Rossby numbers, the skewness of the axial vorticity is expected to diminish, and hence the effect of the large-scale flows on the efficiency of the heat transfer may be different [24].

In all of the simulations described above, the domains are characterized by $L_x = 1$, with differing values of L_y . Although a systematic exploration of L_x - L_y space is not computationally feasible when both L_x and L_y exceed unity, it is nonetheless of interest to investigate the nature of the solution in the regime $1 < L_x < L_y$, given that, for $1 < L_x = L_y$, the attracting solution will consist of only an LSV. To this end, we simulated the case of $L_x = 2$, $L_y = 4$, with the standard parameter values of $\text{Ek} = 10^{-5}$, $\text{Ra} = 3 \times 10^8$, and $\text{Pr} = 1$. Following a small perturbation to the static state, the general form of the evolution is as described above for the cases with $L_x = 1$, with an initial rapid growth and decline in the kinetic energy, followed by a slow evolution to the large-scale state (as shown, for example, in Fig. 1). In this extended domain, the slow evolution is accordingly lengthened, with equilibration of the kinetic energy not occurring until $t \approx 5 \times 10^4$. Figure 16 shows a snapshot of the depth-averaged axial vorticity in the long-time stationary regime, with one jet and with LSVs embedded in the flanks of the cyclonic component; the LSVs extend essentially across the entire x domain. In comparison with the case of $L_x = 1$, $L_y = 2$, which has the same aspect ratio in the horizontal plane, the flows in the extended domain ($L_x = 2$, $L_y = 4$) are, as expected, much more vigorous, with, for example, $\text{Re}_x = 9870$, in comparison with $\text{Re}_x = 5548$ in the smaller domain; as a result, this leads to a reduction in the heat transfer ($\text{Nu} = 26.9$ in comparison with $\text{Nu} = 28.6$).

In spherical shell models of rotating convection in the molecular envelope of gas giant planets, shallow large-scale vortices have been observed to coexist with deep zonal flows (i.e., flows that extend throughout most of the layer depth) [28]. The persistence of deep LSVs when the horizontal symmetry is broken, together with their coexistence with jets as demonstrated here, is a novel feature in rotating convection. We have used a simplified model that does not permit the saturation of the length scale of the large-scale flows as a result of either the curvature of the boundaries or the latitudinal variation of the rotation background. Our model is therefore unable to provide predictions

for the amplitude and length scale of the large-scale flows in terms of the properties of the convective flows. However, it does offer a valuable proof of concept of the possible coexistence of deep LSVs and jets in rotating convection.

In numerical simulations using periodic boundaries, the dynamics should ideally not be influenced by the choice of the computational domain size. However, large-scale flows can be affected by confinement effects due to the finite domain size, as we find here. Numerical simulations in very wide computational domains are required to study the influence of confinement on the large-scale dynamics, but, unfortunately, such simulations are currently out of reach of 3D models of rapidly rotating convection. The analogy between the large-scale dynamics observed in 3D simulations and in forced 2D turbulence simulations (e.g., Frishman *et al.* [17]) supports the study of the evolution of large-scale flows over long time scales using 2D turbulence models in the wide domains that 3D models cannot explore. It would be particularly interesting to perform 2D simulations with a forcing possessing similar statistical properties to the small-scale convective buoyancy driving found in 3D models. An alternative and complementary approach to 3D direct numerical simulations, proposed by Julien and Knobloch [29], is to model rotating Rayleigh-Bénard convection using an asymptotically reduced set of PDEs valid in the limit of small Rossby number. Certain results obtained with the reduced system are found to be in excellent agreement with 3D numerical simulations of the full equations run at moderately low Ekman and Rossby numbers in the absence of domain horizontal anisotropy [18,24]. The reduced model, which assumes the dominance of the planetary vorticity over the planetary vorticity, is, however, unable to capture any asymmetry between cyclonic and anticyclonic motions. Interestingly, the effect of domain horizontal anisotropy on the formation of large-scale flows in rotating Rayleigh-Bénard convection has been recently investigated with the reduced system by Julien *et al.* [30]. They find that unidirectional jets with embedded persistent LSVs form for aspect ratios L_y/L_x in excess of 1.1, in good agreement with our 3D numerical results. The asymptotically reduced system of Julien *et al.* [30] would therefore be a valuable tool to study the confinement effects of the large-scale dynamics in the limit of small Rossby number.

An alternative method of introducing horizontal anisotropy in a Cartesian domain is to tilt the rotation axis with respect to the direction of gravity. This configuration would arguably offer a greater challenge for the formation of the LSVs because of the production of horizontal flows with a vertical shear due to the tilt of the convective rolls [31]. This configuration was investigated by Chan and Mayr [14] for compressible convection. They observed the formation of LSVs in cases with both a small tilt of the rotation axis (corresponding to a latitude of 67.5°) as well as a large tilt (a latitude of 22.5°). The rotation rate required for the appearance of the LSVs is higher at low latitudes. The role played by compressibility in the formation of the LSVs is difficult to assess in this previous work; it would therefore be of great interest to investigate the evolution of the LSVs in Boussinesq convection with a tilted rotation axis.

The collective action of LSVs and rotating convection can produce coherent large-scale magnetic fields at moderate magnetic Reynolds numbers and low Prandtl numbers (the parameter regime relevant for planetary dynamos) despite the disruptive feedback of magnetic fields on coherent Reynolds stresses [9,10]. This magnetic feedback is known to strongly affect unidirectional flows [32,33]. It would therefore be of considerable interest to investigate whether the jets can promote the dynamo mechanism driven by the LSVs and the convection or whether instead they are completely suppressed by the self-sustained magnetic fields. We shall pursue this question in a forthcoming study.

ACKNOWLEDGMENTS

C.G. was supported by the Natural Environment Research Council under Grant No. NE/M017893/1. Computations were performed using the ARCHER UK National Supercomputing Service (www.archer.ac.uk); the facilities of N8 HPC Centre of Excellence, provided and funded by the N8 consortium and EPSRC (Grant No. EP/K000225/1) and coordinated by the Universities of Leeds and Manchester; and ARC1, part of the High Performance Computing facilities at the

University of Leeds. We are grateful to Keith Julien and coauthors for sharing with us the preprint of their recent paper. We thank the two anonymous referees for suggestions that have improved the paper.

-
- [1] C. A. Jones, Thermal and compositional convection in the outer core, in *Treatise on Geophysics*, edited by G. Schubert (Elsevier, Amsterdam, 2007), pp. 131–185.
 - [2] J. M. Aurnou, M. A. Calkins, J. S. Cheng, K. Julien, E. M. King, D. Nieves, K. M. Soderlund, and S. Stellmach, Rotating convective turbulence in Earth and planetary cores, *Phys. Earth Planet. Int.* **246**, 52 (2015).
 - [3] M. Heimpel, J. Aurnou, and J. Wicht, Simulation of equatorial and high-latitude jets on Jupiter in a deep convection model, *Nature (London)* **438**, 193 (2005).
 - [4] P. L. Read, T. N. L. Jacoby, P. H. T. Rogberg, R. D. Wordsworth, Y. H. Yamazaki, K. Miki-Yamazaki, R. M. B. Young, J. Sommeria, H. Didelle, and S. Viboud, An experimental study of multiple zonal jet formation in rotating, thermally driven convective flows on a topographic beta-plane, *Phys. Fluids* **27**, 085111 (2015).
 - [5] C. C. Porco, R. A. West, A. McEwen, A. D. Del Genio, A. P. Ingersoll, P. Thomas, S. Squyres, L. Dones, C. D. Murray, T. V. Johnson, J. A. Burns, A. Brahic, G. Neukum, J. Veverka, J. M. Barbara, T. Denk, M. Evans, J. J. Ferrier, P. Geissler, P. Helfenstein, T. Roatsch, H. Throop, M. Tiscareno, and A. R. Vasavada, Cassini imaging of Jupiter’s atmosphere, satellites, and rings, *Science* **299**, 1541 (2003).
 - [6] U. A. Dyudina, A. P. Ingersoll, S. P. Ewald, A. R. Vasavada, R. A. West, A. D. Del Genio, J. M. Barbara, C. C. Porco, R. K. Achterberg, F. M. Flasar, A. A. Simon-Miller, and L. N. Fletcher, Dynamics of Saturn’s south polar vortex, *Science* **319**, 1801 (2008).
 - [7] C. Guervilly, D. W. Hughes, and C. A. Jones, Large-scale vortices in rapidly rotating Rayleigh-Bénard convection, *J. Fluid Mech.* **758**, 407 (2014).
 - [8] R. K. Yadav, T. Gastine, U. R. Christensen, L. D. V. Duarte, and A. Reiners, Effect of shear and magnetic field on the heat-transfer efficiency of convection in rotating spherical shells, *Geophys. J. Int.* **204**, 1120 (2016).
 - [9] C. Guervilly, D. W. Hughes, and C. A. Jones, Generation of magnetic fields by large-scale vortices in rotating convection, *Phys. Rev. E* **91**, 041001 (2015).
 - [10] C. Guervilly, D. W. Hughes, and C. A. Jones, Large-scale-vortex dynamos in planar rotating convection, *J. Fluid Mech.* **815**, 333 (2017).
 - [11] P. A. Gilman, Nonlinear dynamics of Boussinesq convection in a deep rotating spherical shell. I, *Geophys. Astrophys. Fluid Dyn.* **8**, 93 (1977).
 - [12] F. H. Busse and L. L. Hood, Differential rotation driven by convection in a rapidly rotating annulus, *Geophys. Astrophys. Fluid Dyn.* **21**, 59 (1982).
 - [13] K. L. Chan, Rotating convection in f-boxes: Faster rotation, *Astron. Nachr.* **328**, 1059 (2007).
 - [14] K. L. Chan and H. G. Mayr, Numerical simulation of convectively generated vortices: Application to the Jovian planets, *Earth Plan. Sci. Lett.* **371**, 212 (2013).
 - [15] A. M. Rubio, K. Julien, E. Knobloch, and J. B. Weiss, Upscale Energy Transfer in Three-Dimensional Rapidly Rotating Turbulent Convection, *Phys. Rev. Lett.* **112**, 144501 (2014).
 - [16] M. E. Maltrud and G. K. Vallis, Energy spectra and coherent structures in forced two-dimensional and beta-plane turbulence, *J. Fluid Mech.* **228**, 321 (1991).
 - [17] A. Frishman, J. Laurie, and G. Falkovich, Jets or vortices—What flows are generated by an inverse turbulent cascade? *Phys. Rev. Fluids* **2**, 032602 (2017).
 - [18] S. Stellmach, M. Lischper, K. Julien, G. Vasil, J. S. Cheng, A. Ribeiro, E. M. King, and J. M. Aurnou, Approaching the Asymptotic Regime of Rapidly Rotating Convection: Boundary Layers Versus Interior Dynamics, *Phys. Rev. Lett.* **113**, 254501 (2014).

- [19] R. P. J. Kunnen, R. Ostilla-Mónico, E. P. van der Poel, R. Verzicco, and D. Lohse, Transition to geostrophic convection: The role of the boundary conditions, *J. Fluid Mech.* **799**, 413 (2016).
- [20] F. Cattaneo, T. Emonet, and N. O. Weiss, On the interaction between convection and magnetic fields, *Astrophys. J.* **588**, 1183 (2003).
- [21] S. Chandrasekhar, *Hydrodynamic and Hydromagnetic Stability* (Clarendon, Oxford, 1961).
- [22] P. J. Käpylä, M. J. Mantere, and T. Hackman, Starspots due to large-scale vortices in rotating turbulent convection, *Astrophys. J.* **742**, 34 (2011).
- [23] B. Favier, L. J. Silvers, and M. R. E. Proctor, Inverse cascade and symmetry breaking in rapidly-rotating Boussinesq convection, *Phys. Fluids* **26**, 096605 (2014).
- [24] K. Julien, A. M. Rubio, I. Grooms, and E. Knobloch, Statistical and physical balances in low Rossby number Rayleigh–Bénard convection, *Geophys. Astrophys. Fluid Dyn.* **106**, 392 (2012).
- [25] F. Bouchet and E. Simonnet, Random Changes of Flow Topology in Two-Dimensional and Geophysical Turbulence, *Phys. Rev. Lett.* **102**, 094504 (2009).
- [26] R. M. Clever and F. H. Busse, Three-dimensional convection in a horizontal fluid layer subjected to a constant shear, *J. Fluid Mech.* **234**, 511 (1992).
- [27] J. von Hardenberg, D. Goluskin, A. Provenzale, and E. A. Spiegel, Generation of Large-Scale winds in Horizontally Anisotropic Convection, *Phys. Rev. Lett.* **115**, 134501 (2015).
- [28] M. Heimpel, T. Gastine, and J. Wicht, Simulation of deep-seated zonal jets and shallow vortices in gas giant atmospheres, *Nature Geosci.* **9**, 19 (2016).
- [29] K. Julien and E. Knobloch, Reduced models for fluid flows with strong constraints, *J. Math. Phys.* **48**, 065405 (2007).
- [30] K. Julien, E. Knobloch, and M. Plumley, Impact of domain anisotropy on the inverse cascade in geostrophic turbulent convection, *J. Fluid Mech.* [arXiv:1711.01685](https://arxiv.org/abs/1711.01685).
- [31] D. H. Hathaway and R. C. J. Somerville, Three-dimensional simulations of convection in layers with tilted rotation vectors, *J. Fluid Mech.* **126**, 75 (1983).
- [32] J. Aubert, Steady zonal flows in spherical shell dynamos, *J. Fluid Mech.* **542**, 53 (2005).
- [33] S. M. Tobias, P. H. Diamond, and D. W. Hughes, β -plane magnetohydrodynamic turbulence in the solar tachocline, *Astrophys. J. Lett.* **667**, L113 (2007).